

The Standard Atmosphere

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Some definitions

Absolute altitude

Geometric altitude

Geopotential altitude

Some physics – The hydrostatic equation

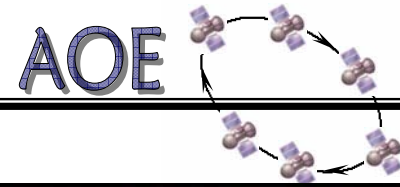
Construction of the standard atmosphere – Variation of p , T and ρ with altitude

Definitions of pressure, density, and temperature altitudes

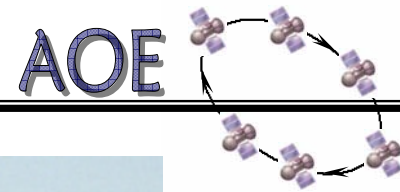
Mysterious Fishy-Looking Thingy



Tourist Transportation?



Upside down?



NASA Dryden flight Research Center Photo Collection
<http://www.dfrc.nasa.gov/gallery/photo/index.html>
NASA Photo: EC03-0154-3 Date: June 9, 2003 Photo By: Tom Tschida

The remotely-piloted Altair unmanned aerial vehicle (UAV) took to the air on its first checkout flight on June 9, 2003 at El Mirage, California.

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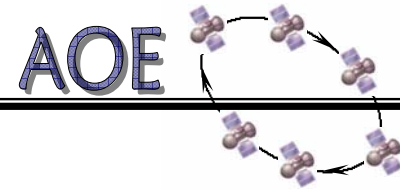
Construction of the standard atmosphere – Variation of p , T and ρ with altitude

Definitions of pressure, density, and temperature altitudes

Altitude (Six Different Flavors)

- Geometric altitude
 h_G : altitude measured from sea level
- Absolute altitude
 h_a : altitude measured from center of the earth
 $h_a = h_G + r$, where r is the radius of the earth
- Local acceleration due to gravity depends on altitude
 $g = g_0 (r/h_a)^2 = g_0 (r/[r+h_G])^2$

This variation has an effect on the pressure



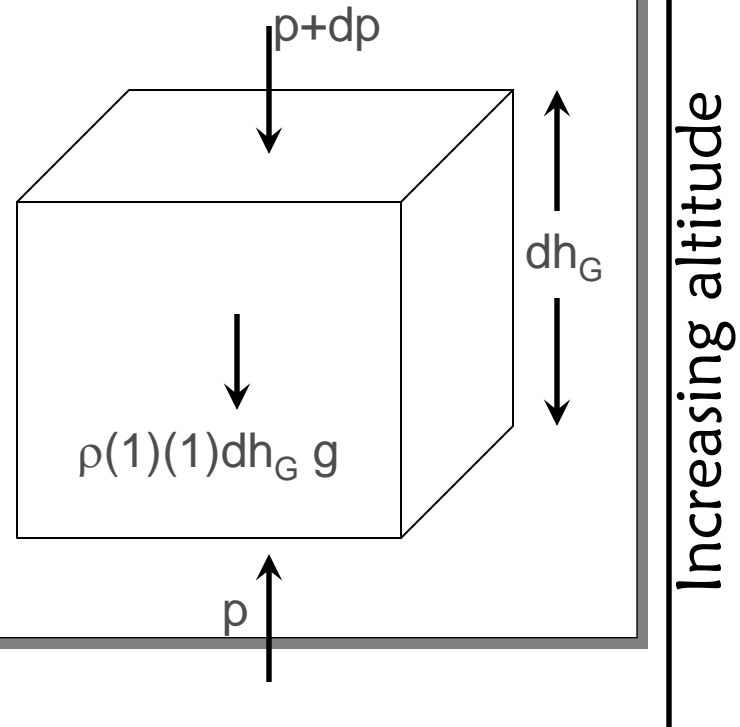
Hydrostatic Equation

- Compute the balance of forces acting on an element of fluid at rest.
- The element of fluid has unit width in both horizontal directions and height dh_G in the vertical direction
- Newton's law: sum of forces = zero

$$p = p + dp + \rho g dh_G$$

$$dp = - \rho g dh_G$$

- The hydrostatic equation applies to any fluid of density ρ
- The H.E. is a differential equation:
 $\rho, \rho,$ and g depend on h_G
- First step: assume $g = g_0$



Integrating the Hydrostatic Eq.

- Making the assumption that $g = g_0$, the H.E. becomes

$$dp = -\rho g_0 dh$$

- Note that we have replaced h_G with h
Since g is not really g_0 , we use h that's not really h_G
The variable h is called the **geopotential altitude**

- What is the relationship between the geopotential and geometric altitudes?

$$dp = -\rho g dh_G \quad \text{and} \quad dp = -\rho g_0 dh$$

Should both be true for the same differential change in pressure, so divide them:

$$1 = (g_0/g)(dh/dh_G) \quad \Rightarrow \quad dh = (g/g_0) dh_G$$

Geopotential and Geometric h

- We assumed $g = g_0$, and defined the geopotential altitude h , based on that assumption, leading to

$$dh = (g/g_0) dh_G$$

- Need to recall that gravity depends on altitude:

$$g = g_0(r/h_a)^2 = g_0 (r/[r+h_G])^2$$

- Thus

$$dh = (g/g_0) dh_G = (r/[r+h_G])^2 dh_G \quad \text{[Exercise] and hence}$$

$$h = r h_G / (r+h_G)$$

Finally, defining the Standard Atmosphere

- We want to know $p(h)$, $T(h)$, and $\rho(h)$
- The fundamental idea behind the standard atmosphere is a *defined* variation of temperature $T = T(h)$
- The figure to the right defines the temperature variation
 - Isothermal and gradient regions
- Most aircraft fly below 20 km, but balloons, sounding rockets, and launch vehicles traverse the entire range

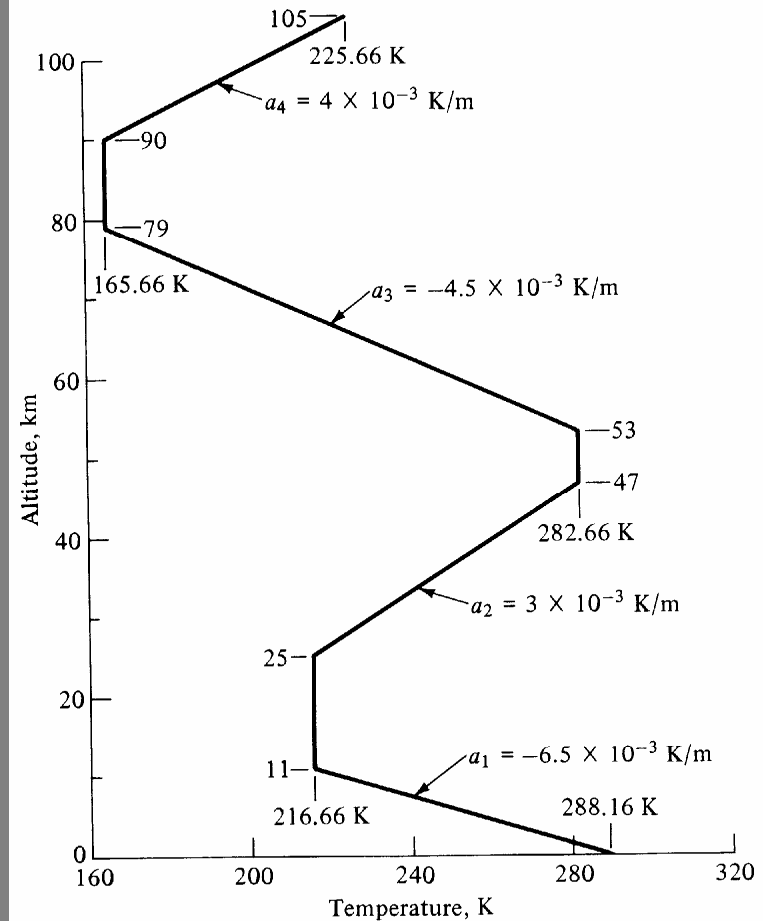


Figure 3.3 Temperature distribution in the standard atmosphere.

Integrating Pressure

- Given the definition of $T(h)$, revisit the equation:
$$dp = - \rho g_0 dh$$
- Remember the equation of state?
$$p = \rho RT$$
- Divide the hydrostatic equation by the equation of state:
$$dp/p = - g_0 dh/(RT)$$
- In the isothermal regions, T is constant, so **[Exercise]**
$$p/p_1 = e^{-[g_0/(RT)](h-h_1)}$$

$$\rho/\rho_1 = e^{-[g_0/(RT)](h-h_1)}$$
- In the gradient regions, T depends on h , so the integration is more difficult

Integrating Pressure (2)

- Begin with:
$$dp/p = -g_0 dh/(RT)$$
- In the gradient regions, T depends linearly on h
$$(T-T_1)/(h-h_1) = dT/dh = a \quad (a = \text{lapse rate})$$
- So,
$$dh = dT/a \quad \Rightarrow \quad dp/p = -g_0 dT/(aRT)$$
- Integrating **[Exercise]** leads to
$$p/p_1 = (T/T_1)^{-g_0/(aR)}$$
$$\rho/\rho_1 = (T/T_1)^{-[g_0/(aR)+1]}$$
- Since $T=T(h)$, we can write $p=p(h)$ and $\rho=\rho(h)$
- And it all starts at sea level, where $h=h_G=0$

At Sea Level

- At s.l., $h=h_G=0$
 - $p_s = 1.01325 \times 10^5 \text{ N/m}^2 = 2116.2 \text{ lb/ft}^2$
 - $\rho_s = 1.2250 \text{ kg/m}^3 = 0.002377 \text{ slug/ft}^3$
 - $T_s = 288.16 \text{ K} = 518.69^\circ\text{R}$
- From sea level, we can calculate the rest
- The calculated values are available in Appendices A & B in the text, in SI and English units, respectively

Summary of Standard Atmosphere

- Temperature distribution defined by Fig. 3.4
- Equation of state relates pressure, density and temperature
- Integration of hydrostatic equation and use of equation of state leads to:

Isothermal

$$p/p_1 = e^{-[g_0/(RT)](h-h_1)}$$

$$\rho/\rho_1 = e^{-[g_0/(RT)](h-h_1)}$$

$$T = \text{constant}$$

Gradient

$$p/p_1 = (T/T_1)^{-g_0/(aR)}$$

$$\rho/\rho_1 = (T/T_1)^{-[g_0/(aR)+1]}$$

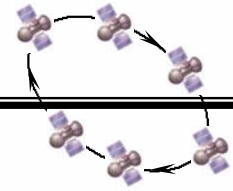
$$T = T_1 + a(h-h_1)$$

- See Appendices A & B in text

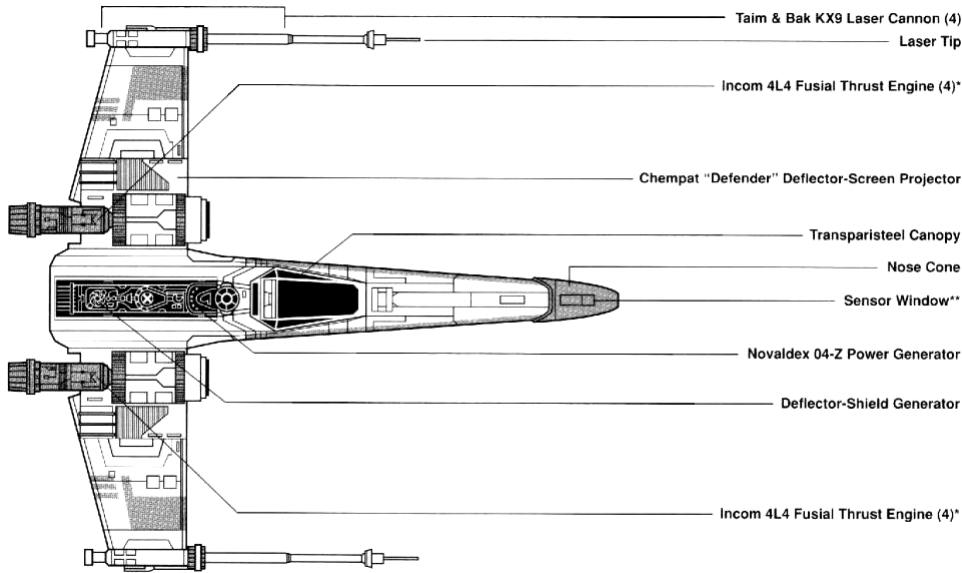
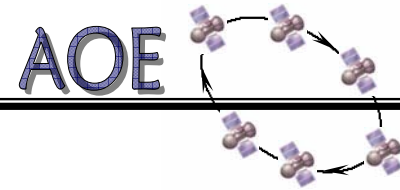
Who Is This Guy and What Is He Doing?



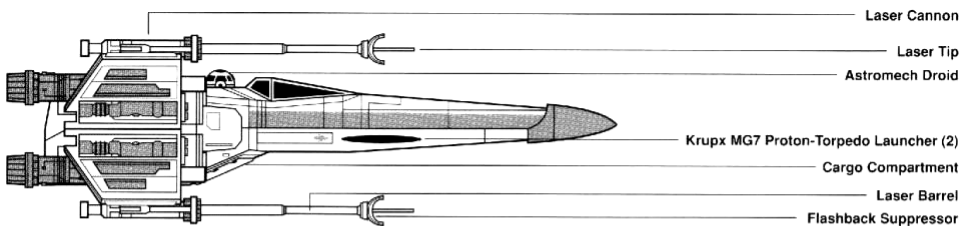
From Where to Here? AOE



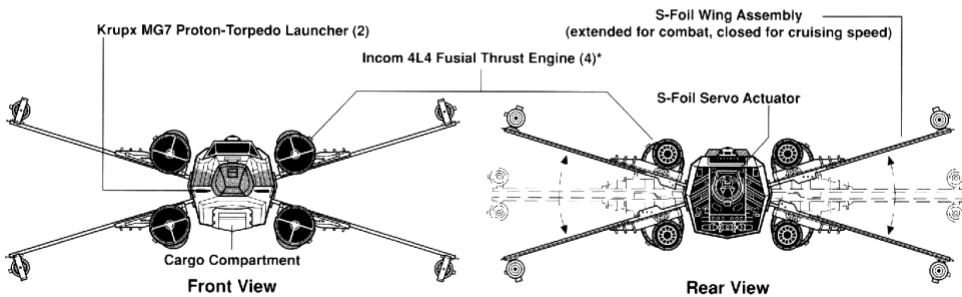
Is It Real?



Top View



Side View



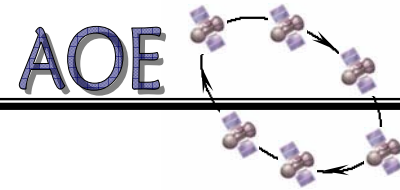
*Alternate configurations may use Incom 4j.4 fusial thrust engines.

** Houses Carbanti Transceiver Package with Fabritech AN6-5d "lock track" full-spectrum transceiver, Mellihat "Multi Imager" dedicated energy receptor, Taha Iro electro-photo receptor and Fabritech AN6 3.6 sensor computer. Alternate configuration typically combines long-range Fabritech AN6-5d units with long-range PTDA #9A-9r unit and short-range PTAG #PG-7u unit.

Motion Sickness Generator



Example 3.1



Calculate the standard atmosphere values of T , p , and ρ at a geopotential altitude of 14 km.

From Fig. 3.4, $T = 216.66$ K (isothermal region)

Begin at sea level and compute values at the “corner”:

Gradient region from $h = 0$ to $h = 11.0$ km,

with lapse rate $a = -6.5$ K/km

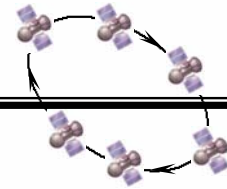
So $p = p_1(T/T_1)^{-g_0/(aR)}$ and $\rho = \rho_1(T/T_1)^{-[g_0/(aR)+1]}$

$p_1 = p_s = 1.01325 \times 10^5$ N/m² and $\rho_1 = \rho_s = 1.2250$ kg/m³

$T_1 = T_s = 288.16$ K

$p = 2.26 \times 10^4$ N/m² and $\rho = 0.367$ kg/m³

at $h = 11.0$ km



From 11.0 km to 14.0 km is an isothermal region, so

$$p = p_1 e^{-[g_0/(RT)](h-h_1)} \quad \text{and} \quad \rho = \rho_1 e^{-[g_0/(RT)](h-h_1)}$$

Here the subscript “1” refers to the values at $h=11.0$ km

Doing the calculations leads to

$$p = 1.41 \times 10^4 \text{ N/m}^2 \quad \text{and} \quad \rho = 0.23 \text{ kg/m}^3$$

at $h = 14.0$ km

Three More Altitudes

- Pressure, Temperature and Density Altitudes
- Remember Geometric, Absolute, and Geopotential Altitudes
- Suppose you're flying and you have a set of instruments that can measure Pressure, Temperature and Density
- You could look them up in the standard atmosphere tables, right?
- Would they all agree? Not likely.

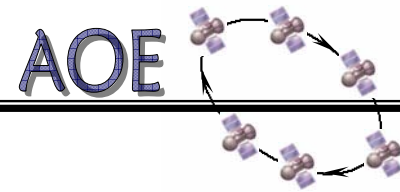
1976 Standard Atmosphere

(from <http://www.digitaldutch.com/atmoscalc/tableoptions1.htm>)

Altitude [km]	Temperature [Kelvin]	Pressure [pascal]	Density [kg/m ³]
0	288.15	101325	1.225
1	281.65	89874.5705	1.1116
2	275.15	79495.2155	1.0065
3	268.65	70108.5447	0.9091
4	262.15	61640.2353	0.8191
5	255.65	54019.9121	0.7361

- Suppose measured pressure is $6.16 \times 10^4 \text{ N/m}^2$
 - Then the pressure altitude is $\approx 4 \text{ km}$
- Suppose the measured temperature is 269 K
 - Then the temperature altitude is $\approx 3 \text{ km}$
- What is the geometric altitude?

Example 3.2



If an airplane is flying at an altitude where the actual pressure and temperature are

$$4.72 \times 10^4 \text{ N/m}^2 \quad \text{and} \quad 255.7 \text{ K}$$

respectively, what are the pressure, temperature and density altitudes?

Use Appendix A to find that

$$\text{pressure altitude} = 6 \text{ km}$$

$$\text{temperature altitude} = 5 \text{ km (or 38.2 or 59.5 km)}$$

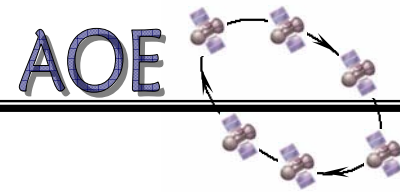
Why?

Use equation of state to find that the density is

$$\rho = 0.643 \text{ kg/m}^3$$

And then use Appendix A to find that

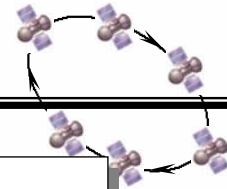
$$\text{density altitude} = 6.24 \text{ km}$$



- Standard atmosphere is defined in order to relate flight tests, wind tunnel results, and general airplane design and performance to a common reference
- Temperature variation is defined
 - Isothermal and gradient regions
- Hydrostatic equation is applied
$$dp = - \rho g dh_G$$
- Pressure and density variations are derived
- There are six different flavors of altitude:
 - Geometric, absolute, geopotential
 - Pressure, temperature, density

Additional Topics

- Geometric vs geopotential altitudes
- Linear interpolation
- Geopotential altitude is only used as a convenience for computing the standard atmosphere tables
- Always use the geometric altitude column of the tables when referring to pressure, density, and temperature altitudes



- Usually looking up numbers in tables requires *linear interpolation*
 - Select the two rows that “bracket” the given value
 - Use a straight-line approximation between the values in the two rows

h [km]	T [K]	P [N/m ²]	ρ [kg/m ³]
1	281.65	89874	1.1116
2	275.15	79495	1.0065

Example: you measure pressure at

$$p^* = 8.0 \times 10^4 \text{ N/m}^2$$

Then the pressure altitude is found by

$$h^* = h_1 + (h_2 - h_1)(p^* - p_1)/(p_2 - p_1) \\ = 1.9514 \text{ km}$$

